Methane emission by plant communities in an alpine meadow on the Qinghai-Tibetan Plateau: a new experimental study of alpine meadows and oat pasture

Sheping Wang1,2, Xiaoxia Yang1,2, Xingwu Lin3, Yigang Hu1,2, Cajun Luo1,2, Guangping Xu1,2, Zhenhua Zhang1,2, Ailing Su1,2, Xiaofen Chang1,2, Zengguo Chao1,2 and Jichuang Duan1,2

1Key Laboratory of Adaptation and Evolution of Plateau Biota, Northwest Institute of Plateau Biology, Chinese Academy of Sciences, Xining 810008, People’s Republic of China
2Graduate University of Chinese Academy of Sciences, Beijing 100049, People’s Republic of China
3State Key Laboratory of Soil and Sustainable Agriculture, Institute of Soil Science in Nanjing, Chinese Academy of Sciences, Nanjing 210008, People’s Republic of China
*Author for correspondence (wangship2008@yahoo.cn).

Recently, plant-derived methane (CH4) emission has been questioned because limited evidence of the chemical mechanism has been identified to account for the process. We conducted an experiment with four treatments during the growing seasons of 2007 and 2008 to examine the hypothesis that abiotic (i.e. soil moisture and soil temperature) rather than biotic (i.e. alpine vegetation) factors resulted in the difference in CH4 consumption between plots with vegetation and plots with bare soil.

2. MATERIAL AND METHODS

The study site was the same as the alpine Kobresia meadow studied by Cao et al. (2008) during the growing seasons of 2007 and 2008. A completely randomized design was employed, with four replicate plots of each of four treatments as follows: (i) native, natural alpine meadow; (ii) naturally restored alpine meadow eight years after cultivation in the 1960s; (iii) bare soil with roots removed from 0–20 cm soil depth in May 2007; and (iv) annual oat sown with 600 kg seeds per hectare in mid-June 2007 and by the end of May 2008. Each plot (4 × 4.5 m) was separated by a 2 m buffer zone. The total rainfall was 352 and 290 mm from June to September in 2007 and 2008, respectively.

Fluxes of CH4 were measured weekly inside opaque, static, stainless steel chambers using the methods of flux calculation described by Ma et al. (2006). CH4 concentrations of gas samples were analysed by gas chromatography (HP Series 4890D, Hewlett Packard, USA) within 24 h. The fluxes of CH4 between 9.00 a.m. and 11.00 a.m. local time were used to represent 1 day’s average flux as described by Cao et al. (2008).

During each gas-sampling occasion in 2007, soil temperature was measured using digital thermometers in situ at 5 cm depth in all plots. The volumetric soil moisture (%) at 5 cm depth was measured during both 2007 and 2008 using time domain reflectometry (CS615) for each plot to calculate the water-filled pore space (WFPS):

\[
WFPS = \frac{\text{volumetric soil moisture}}{(1 - BD/DD) + 0.01}
\]

where BD is the soil bulk density and PD is the soil particle density.

A general linear model repeated-measures define factors procedure (SPSS 12.0, SPSS Inc., Chicago, IL, USA) was used to assess the significance of the impacts of experimental year, sampling day, treatment and their interaction on soil water content, soil temperature and CH4 fluxes. The significant differences between treatments were assessed by one-way ANOVA and least significant difference. The Pearson’s correlations were calculated between soil temperature and WFPS and CH4 fluxes. All significances mentioned in the text were at 0.05 level.

3. RESULTS

(a) Environmental changes

Soil WFPS was affected significantly by treatment, sampling date, year and their interaction (table 1). The average soil water content for both years in the bare soil plots (31.8 ± 1.0%) was significantly lower (by approx. 24%) than in the native, natural alpine meadow and naturally restored alpine meadow plots, whereas higher (by 11%) than in the oat plots. The average soil temperature at 5 cm soil depth was not

Table 1. Soil WFPS and methane (CH$_4$) consumption rates from repeated-measures ANOVA using year and sampling date as repeated measures (between-subjects).

<table>
<thead>
<tr>
<th>model</th>
<th>WFPS</th>
<th>CH$_4$ consumption rate</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>MS</td>
<td>$F$</td>
</tr>
<tr>
<td>treatment (T)</td>
<td>1381.57</td>
<td>34.42</td>
</tr>
<tr>
<td>year (Y)</td>
<td>10 456.5</td>
<td>467.49</td>
</tr>
<tr>
<td>Y $\times$ T</td>
<td>297.64</td>
<td>13.307</td>
</tr>
<tr>
<td>date (D)</td>
<td>1076.46</td>
<td>106.33</td>
</tr>
<tr>
<td>D $\times$ T</td>
<td>24</td>
<td>2.371</td>
</tr>
<tr>
<td>Y $\times$ D</td>
<td>557.82</td>
<td>47.421</td>
</tr>
<tr>
<td>Y $\times$ D $\times$ T</td>
<td>13.89</td>
<td>1.18</td>
</tr>
</tbody>
</table>

Figure 1. Relationships between methane (CH$_4$) consumption rate and (a) soil temperature in 2007 and (b) soil WFPS for both years significantly different between all treatments during the study period in 2007 (data not shown).

(b) Methane consumption rate
The effect of different land-use (treatment effect) on CH$_4$ consumption rate for every sample date was not significant for either year (table 1). However, multiple comparison analysis between mean values for each sampling period showed significant differences in CH$_4$ consumption rate between native, natural alpine meadow and oat plots in 2007, and between native, natural alpine meadow and oat and bare soil plots in 2008 (electronic supplementary material). CH$_4$ consumption rate in 2007 increased with an increase in soil temperature at 5 cm depth (figure 1a). However, increasing WFPS significantly decreased CH$_4$ consumption (figure 1b).

(c) Methane emission by alpine communities
Our study showed that bare soil plots were a net sink for atmospheric CH$_4$ with an average of approximately 40.7 $\mu$g CH$_4$ m$^{-2}$ h$^{-1}$ (range: $-14.8 \sim -63.7$) in 2007 and 52.5 $\mu$g CH$_4$ m$^{-2}$ h$^{-1}$ ($-17 \sim -79$) in 2008 during the study periods. Our calculation assumed that the CH$_4$ emission rate by plant communities was the difference between plots treated with vegetation and bare soil (Cao et al. 2008) (which we have called apparent emission rate by plants). The average apparent CH$_4$ emission rates were 15.0 $\mu$g CH$_4$ m$^{-2}$ h$^{-1}$ (9.9 $\sim$ 20.2) and 5.1 $\mu$g CH$_4$ m$^{-2}$ h$^{-1}$ (0.3 $\sim$ 9.9), with great seasonal variations of $-65.0 \sim 75.0$ $\mu$g CH$_4$ m$^{-2}$ h$^{-1}$ in 2007 and $-48.0 \sim 72.0$ $\mu$g CH$_4$ m$^{-2}$ h$^{-1}$ in 2008 for the native, natural alpine meadow community and the naturally restored alpine meadow community, respectively. In contrast, annual oat vegetation apparently consumed atmospheric methane at an average rate of 4.8 $\mu$g CH$_4$ m$^{-2}$ h$^{-1}$ ($-0.8 \sim -8.7$ $\mu$g CH$_4$ m$^{-2}$ h$^{-1}$), with great seasonal variations of $-5.8 \sim 13.4$ $\mu$g CH$_4$ m$^{-2}$ h$^{-1}$ in 2007 and $-3.0 \sim 27.0$ $\mu$g CH$_4$ m$^{-2}$ h$^{-1}$ in 2008 (figure 2).

4. DISCUSSION

(a) Methane emissions by plants

Our results appear to indicate possible CH$_4$ production by alpine meadow vegetation because the two types of meadow vegetation reduced CH$_4$ consumption compared with bare soil. In contrast, the oat pasture increased CH$_4$ consumption compared with bare soil. These results seem to support the apparent conclusion that the intact Kobresia meadow emitted CH$_4$ as reported by Cao et al. (2008). However, the response of CH$_4$ consumption is very sensitive to changes in soil moisture and temperature in the field. In our study, we calculated a regression equation between CH$_4$ consumption rate and soil WFPS and soil temperature in 2007, which was:

$$y = -88.286 + 1.135 \text{WFPS} - 0.164 T$$

where $T$ is the soil temperature at 5 cm depth. Thus, native, natural meadow plots could consume more CH$_4$ (approx. 11.8 µg CH$_4$ m$^{-2}$ h$^{-1}$) through calibration based on the same soil WFPS and temperature as bare soil plots. This effect would explain the original difference of the CH$_4$ consumption rate between native, natural alpine meadow ($-30.8$ µg CH$_4$ m$^{-2}$ h$^{-1}$) and bare soil ($-40.7$ µg CH$_4$ m$^{-2}$ h$^{-1}$) in 2007. Similar results were observed in 2008 and for other treatments in our study. Therefore, our results strongly suggest that the apparent CH$_4$ production by vegetation, when calculated in comparison with bare soil in some previous studies, might represent differences in soil temperature and WFPS and not the true vegetation sources of CH$_4$.

(b) Effect of soil temperature and moisture on CH$_4$ consumption

Removal of roots from soil often alters its physico-chemical characteristics, which are of critical importance for CH$_4$ uptake (Smith et al. 2000). Soil temperature measured on nine occasions in 2007 varied between 7 and 24°C and was linearly correlated with soil WFPS ($r = -0.45$, $p < 0.01$) in all treatments. Our study showed a clear positive relationship between CH$_4$ consumption rate and soil temperature for all treatments (figure 2a). Probably, when fewer soil pores are water-filled, more atmospheric CH$_4$ could diffuse into the soil and reach methanotrophic micro-organisms, which might respond positively to the temperature increase (Pearce & Clymo 2001; Zhuang et al. 2007).

In our study, CH$_4$ consumption appeared to increase linearly with decreases in soil moisture (figure 2b). We found that the response of CH$_4$ consumption to soil moisture was greater in 2008 than in 2007 (the slopes of the regression equations between CH$_4$ consumption rate and WFPS were 1.04 and 2.27 in 2007 and 2008, respectively). This may have resulted from more drought in 2008 than in 2007, which would limit the diffusive transport of methane through the soil gas phase when soil moisture is high (King 1997; Castaldi & Fierro 2005). In these soils, where gas diffusion represents the main controlling factor of CH$_4$ oxidation, soil water content is of critical importance in determining the potential of the ecosystem to be a CH$_4$ sink (Striegl 1993).

(c) Effect of solar radiation on CH$_4$ emission

Many studies show that solar radiation stimulates some CH$_4$ emission from plant foliage (Keppeler et al. 2006, 2008; McLeod et al. 2008; Vigano et al. 2008; Messenger et al. 2009), whereas the study of Cao et al. (2008) reported methane emissions from whole plants in plots when compared with bare soil. Our method showed, however, that the CH$_4$ production was different from that reported by Cao et al. (2008). In this study, we observed a new explanation for the apparent methane emissions reported by Cao et al. (2008). We used closed, opaque chambers (i.e. without solar radiation available to plants), whereas Cao et al. (2008) used transparent chambers shaded with white plastic (i.e. with some solar radiation available to the plants). The difference between our results and those of Cao et al. (2008) may partially derive from the different experimental methods used to assess CH$_4$ emissions. However, we also detected that apparent CH$_4$ emissions may arise from treatments owing to changes in soil temperature and WFPS and do not represent a true vegetation source.

Therefore, the question of aerobic methane production from vegetation in the Qinghai-Tibetan Plateau still remains open. Further studies should evaluate the effects of soil conditions on CH$_4$ emission by plant communities and the role of solar radiation, which was excluded from our study of the alpine meadow.

This research was funded by the Knowledge Innovation Programs (KZCX2-XB2-06-01, KSCX2-YW-N-040) and the ‘100-Talent Program’ of Chinese Academy of Sciences and Chinese National Natural Science Foundation Commission (30871824).


